

1 SO₂ over central China: Measurements, numerical simulations 2 and the tropospheric sulfur budget

3 Hao He,¹ Can Li,^{2,3} Christopher P. Loughner,^{2,3} Zhanqing Li,^{1,3,4} Nickolay A. Krotkov,²
4 Kai Yang,^{1,2} Lei Wang,⁴ Youfei Zheng,⁴ Xiangdong Bao,⁵ Guoqiang Zhao,⁵
5 and Russell R. Dickerson^{1,3,6}

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7 [1] SO₂ in central China was measured in situ from an aircraft and remotely using the
8 Ozone Monitoring Instrument (OMI) from the Aura satellite; results were used to develop a
9 numerical tool for evaluating the tropospheric sulfur budget - sources, sinks, transformation
10 and transport. In April 2008, measured ambient SO₂ concentrations decreased from
11 ~7 ppbv near the surface to ~1 ppbv at 1800 m altitude (an effective scale height of
12 ~800 m), but distinct SO₂ plumes were observed between 1800 and 4500 m, the aircraft's
13 ceiling. These free tropospheric plumes play a major role in the export of SO₂ and in the
14 accuracy of OMI retrievals. The mean SO₂ column contents from aircraft measurements
15 (0.73 DU, Dobson Units) and operational OMI SO₂ products (0.63 ± 0.26 DU) were close.
16 The OMI retrievals were well correlated with in situ measurements (r = 0.84), but showed
17 low bias (slope = 0.54). A new OMI retrieval algorithm was tested and showed improved
18 agreement and bias (r = 0.87, slope = 0.86). The Community Multiscale Air Quality
19 (CMAQ) model was used to simulate sulfur chemistry, exhibiting reasonable agreement
20 (r = 0.62, slope = 1.33) with in situ SO₂ columns. The mean CMAQ SO₂ loading over
21 central and eastern China was 54 kT, ~30% more than the estimate from OMI SO₂
22 products, 42 kT. These numerical simulations, constrained by observations, indicate that
23 ~50% (35 to 61%) of the anthropogenic sulfur emissions were transported downwind, and
24 the overall lifetime of tropospheric SO₂ was 38 ± 7 h.

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27 1. Introduction

28 [2] Driven by the rapid economic development in the past
29 decades, the consumption of energy and raw material in
30 China increased dramatically. Coal burning accounts for
31 70% of the total energy consumption in China [CESY,
32 2005], and estimated total anthropogenic sulfur dioxide
33 (SO₂) emissions were ~31.3 Tg in 2008 [Lu et al., 2010].
34 Atmospheric SO₂ is oxidized to form sulfate (SO₄²⁻) aerosols
35 and leads to acid deposition through sulfuric acid (H₂SO₄).
36 The sulfate aerosols can exert influence on weather and cli-
37 mate [Intergovernmental Panel on Climate Change, 2007;
38 Stier et al., 2007], cause visibility impairments [Hand and

Malm, 2007], and pose a hazard to public health [U.S. 39
Environmental Protection Agency, 2004; *He et al.*, 2002; 40
Hu et al., 2010; *Kan et al.*, 2010; *Schlesinger and Cassee*, 41
2003]. These sulfur-compounds can be transported far from 42
the source regions [Dunlea et al., 2009; Prospero et al., 2003; 43
Singh et al., 2009; *van Donkelaar et al.*, 2008]. 44

[3] A number of studies have been conducted to investi- 45
gate the sulfurous pollution in China. Surface observations 46
of SO₂ were made in and near Beijing [C. Li et al., 2007; 47
Sun et al., 2009], Yangtze River Delta (YRD) [Costabile 48
et al., 2006], Pearl River Delta (PRD) [Zhang et al., 2008], 49
and rural areas [Meng et al., 2010]. Aircraft measurements 50
were also performed to study the vertical distribution of SO₂ 51
in the Northeast [Dickerson et al., 2007], South [Wang et al., 52
2008], and East of China [Geng et al., 2009; Xue et al., 53
2010]. Both surface and airborne measurements demon- 54
strated high SO₂ concentrations with large variations in 55
spatial and temporal distributions. For instance, ambient SO₂ 56
measurements in ten background and rural sites revealed 57
concentrations (±standard deviation, σ) of 0.7 ± 0.4 ppbv at 58
Waliguan on Qinghai Plateau and 67.3 ± 31.1 ppbv at Kaili 59
in Southwest China [Meng et al., 2010]; over the PRD, 60
investigators observed 18.5 ppbv SO₂ at 2100 m and up to 61
107.5 ppbv SO₂ within the planetary boundary layer (PBL) 62

¹Department of Atmospheric and Oceanic Science, University of Maryland, College Park, Maryland, USA.

²NASA Goddard Space Flight Center, Greenbelt, Maryland, USA.

³Earth System Science Interdisciplinary Center, University of Maryland, College Park, Maryland, USA.

⁴Nanjing University of Information Science and Technology, Nanjing, China.

⁵Henan Meteorological Bureau, Zhengzhou, China.

⁶Department of Chemistry and Biochemistry, University of Maryland, College Park, Maryland, USA.



Figure 1. Location of Henan province and Y-7 research airplane used for air sampling.

63 during one research flight [Wang *et al.*, 2008]; over north-
64 eastern China 5 ~ 20 ppbv SO₂ in the PBL and <1 ppbv SO₂
65 aloft were observed [Dickerson *et al.*, 2007].

66 [4] Remote sensing of tropospheric SO₂ over China using
67 the OMI instrument has been used to track an individual SO₂
68 plume [Krotkov *et al.*, 2008; Li *et al.*, 2010a] and to identify
69 changes in emission sources [Li *et al.*, 2010b; Witte *et al.*,
70 2009]. OMI SO₂ products showed reasonable agreement
71 with the in situ measurements and estimated SO₂ emission
72 reductions based on a bottom-up approach. Numerical
73 regional air quality models, such as the USEPA CMAQ
74 model, have been employed to simulate the SO₂ chemistry
75 and transport in East Asia [Lin *et al.*, 2008; Liu *et al.*, 2010;
76 Wang *et al.*, 2010a, 2010b]. However, most of the studies
77 focused on the highly industrialized regions of eastern China,
78 and it is crucial to also investigate SO₂ pollution and chem-
79 istry in the developing areas of central and western China.

80 [5] The majority of tropospheric SO₂ is removed by dry
81 deposition or oxidization to form sulfate aerosols (SO₄²⁻).
82 The SO₂ dry deposition velocity has been measured at
83 0.2 ~ 0.4 cm/s in northern China [Clarke *et al.*, 1997;
84 Sorimachi *et al.*, 2003; Sorimachi and Sakamoto, 2007;
85 Wesely and Hicks, 2000]. In the PBL, the lifetime of SO₂ is a
86 few days due to dry deposition alone [Berglen *et al.*, 2004;
87 Chin *et al.*, 1996], and the observed lifetime is greatly
88 decreased by oxidation processes. When lifted to the free
89 troposphere (FT, higher than 2000 m), the atmospheric SO₂
90 has a longer lifetime, and the long-range transport of atmo-
91 spheric SO₂ from China becomes important [Igarashi *et al.*,
92 2006; Kim *et al.*, 2001; Tu *et al.*, 2004]. In the U.S., it was
93 estimated that ~30% of the emitted SO₂ is subsequently
94 removed through dry deposition and ~37% is exported
95 [Shannon and Sisterson, 1992]. During transport, the reac-
96 tive SO₂ is oxidized to form sulfate aerosols and other sul-
97 furous compounds [Calyert *et al.*, 1978; Cox and Penkett,
98 1971; Eggleton and Cox, 1978; Lee *et al.*, 2011]. Sulfate
99 aerosols impact the global radiative balance through direct
100 effects [Haywood and Boucher, 2000; Stier *et al.*, 2007] and
101 indirect effects on clouds [Albrecht, 1989; Twomey, 1977].
102 Therefore, a budget analysis including the long-range trans-
103 port and evolution of tropospheric SO₂ is essential to inves-
104 tigate regional air quality and large-scale climate effects.

105 [6] In April 2008, a joint China-U.S. field experiment was
106 carried out under the East Asian Study of Tropospheric
107 Aerosols and their Impact on Regional Climate (EAST-AIRC)
108 [Li *et al.*, 2011], following the East Asian Study of Tropo-
109 spheric Aerosols, an International Regional Experiment
110 (EAST-AIRE) [Z. Q. Li *et al.*, 2007]. In addition to ground-
111 based campaigns at four locations in southern and northern

China, an airborne campaign on ambient SO₂ was conducted 112
in Henan province, central China (Figure 1). Coalmines and 113
power plants are concentrated in western and southern Henan, 114
and coal burning is ubiquitous for domestic cooking and 115
heating. In 2007, Henan emitted 2.3 Tg SO₂, around 7% of the 116
total emission of China, and the power plants emissions were 117
1.1 Tg SO₂, ranking 3rd in China by province [Lu *et al.*, 2010]. 118
Central China is a major source of sulfur emissions that has not 119
been thoroughly investigated. 120

[7] In this article, we present a study of tropospheric sul- 121
furous pollutants over central China, employing in situ 122
measurements (aircraft campaign), satellite remote sensing 123
(NASA OMI SO₂ products), and numerical simulations (the 124
CMAQ system). Section 2 presents the data set and methods 125
applied. Section 3 describes results from aircraft measure- 126
ments of tropospheric SO₂. In section 4, we compare OMI 127
SO₂ products with in situ measurements. Section 5 describes 128
the set-up of CMAQ system, management of emission 129
inventory, and model modification. In section 6 we discuss 130
the CMAQ results, model evaluation, SO₂ lifetime and 131
transport of sulfur-compounds. Finally we summarize the in 132
situ observations, remotely sensed observations and numer- 133
ical simulations to estimate the fraction of sulfur emitted into 134
the atmosphere then exported from central China to the 135
atmosphere over the Pacific. 136

2. Data and Method 137

[8] For the aircraft campaign, a Y-7 turboprop transport 138
aircraft (the Chinese version of Antonov An-26, Figure 1) 139
was employed as the measurement platform. This airplane 140
with a cruise speed of 400 km/h was based at the Xinzheng 141
International Airport (IATA code: CGO, 34.52°N, 113.84°E) 142
in suburban Zhengzhou, the capitol city of Henan with 143
7 million residents. The aft-facing inlet and the temperature/ 144
relative humidity (T/RH) probe provided by the local Henan 145
Meteorological Bureau (HMB) were installed on a rack to 146
the left of fuselage. HMB also provided GPS and other 147
logistical support. 148

[9] A commercially available Thermo Electron Corpora- 149
tion (TECO) 43C trace level pulsed fluorescence SO₂ ana- 150
lyzer was modified to measure the ambient SO₂ [Luke, 151
1997], with a detection limit enhanced to ~0.3 ppbv for 152
10 s averaging time [Hains, 2007]. The instrument was 153
calibrated with a National Institute of Standard and Tech- 154
nology (NIST) traceable SO₂ gas standard (Scott Marrin 155
INC., Riverside CA, U.S.A.). HMB provided 1 s data of 156
altitude, longitude, latitude, ambient temperature (T) and 157
relative humidity (RH), which were processed to 10 s aver- 158
age data and checked for quality assurance. 159

[10] Flight routes were selected based on the near real- 160
time OMI SO₂ maps (http://so2.gsfc.nasa.gov/pix/daily/0408/china_0408z.html) provided by the NASA OMI SO₂ 161
group. Synoptic circulation patterns were tracked through 162
satellite images, surface analyses and forecasts (<http://web.kma.go.kr/eng/index.jsp>) from the Korea Meteorological 163
Administration (KMA). Flight plans were designed to mea- 164
sure ambient SO₂ over locations with both strong and weak 165
OMI SO₂ signals under different weather conditions. We 166
took off around 1 P.M. local time (0500 UTC), so the spirals 167
were conducted between 1 and 2 P.M. local time, close to 168
the OMI overpass time (1:45 P.M.). The research flights 169
170
171

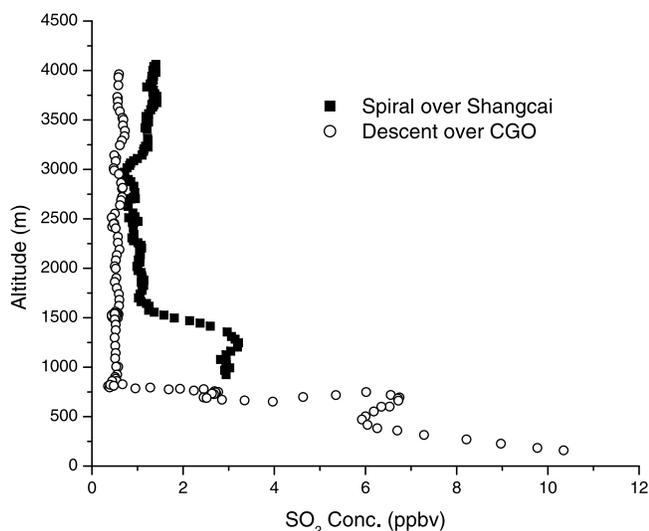


Figure 2. SO₂ Altitude profiles from research flight 04/15/2008. Shangcai (33.25°N, 114.26°E) was a moderate OMI SO₂ signal region, and the descent was over CGO airport (34.52°N, 113.84°E).

172 were confined within the province, and the spirals were
173 restricted to altitudes of 900 ~ 4500 m for safety concerns.
174 We retrieved the information of SO₂ within the PBL (lower
175 than 1000 m, a typical PBL height during spring in China)
176 during the descents into the airport.

177 [11] The backscattered UV radiation measurements from
178 OMI were used to retrieve SO₂ PBL column amounts
179 through the operational Band Residual Difference (BRD)
180 algorithm [Krotkov *et al.*, 2006]. This algorithm was vali-
181 dated over Northeastern China during the EAST-AIRE
182 campaign [Krotkov *et al.*, 2008] and utilized to create oper-
183 ational products. Daily OMI SO₂ and cloud composite ima-
184 ges were available from the NASA SO₂ web site (<http://so2.gsfc.nasa.gov>). In this study, we applied the OMI daily
186 gridded products (SO2L2G, hereafter named OMISO2 pro-
187 ducts, available at http://disc.sci.gsfc.nasa.gov/Aura/data-holdings/OMI/OMISO2g_v003.shtml). These daily products
189 were filtered to remove data with high radiative cloud frac-
190 tion (OMI cloud fraction >0.3) and large solar zenith angle
191 (SZA > 70°).

192 [12] In an effort to improve the detection and quantifica-
193 tion of SO₂ from OMI, we have also utilized an advanced
194 retrieval technique, the iterative spectral fitting (ISF) algo-
195 rithm, previously applied to volcanic clouds [Yang *et al.*,
196 2009a, 2009b, 2010], to take advantage of the large num-
197 ber of spectral measurements available from the hyper-
198 spectral instruments, such as OMI and GOME-2. The ISF
199 algorithm provides less noisy and potentially more accurate
200 column estimates under the diverse range of conditions
201 encountered in global observations, and has been extended
202 to extract the height of a volcanic SO₂ layer in the atmo-
203 sphere [Yang *et al.*, 2009a, 2010]. The ISF products were
204 generated deliberately off-line for 2008 campaign over the
205 East Asia area.

206 [13] We applied the WRF (Weather Research and Fore-
207 casting) - MCIP (Meteorology-Chemistry Interface Proces-
208 sor) - CMAQ system to conduct numerical simulations for
209 our campaign. NASA 2006 Intercontinental Chemical

Transport Experiment Phase-B (INTEX-B) emission inven- 210
tory [Zhang *et al.*, 2009] was utilized to create emission input 211
data for CMAQ. MCIP was developed to correct SO₂ dry 212
deposition velocity, and CMAQ was modified to calculate 213
the flux of pollutants entering and exiting the modeling 214
domain. Hourly outputs of CMAQ simulations were stored 215
and analyzed. Details on the model system are described 216
in section 5. 217

3. Results of Aircraft Campaign 218

[14] During the month-long aircraft campaign for cloud- 219
seeding operations, seven research flights were conducted 220
on April 4, 5, 15, 16, 18, 20, and 22, 2008. These flights 221
were within 150 km of the CGO airport, covering regions 222
with strong OMI SO₂ signal such as Changyuan (34.52°N, 223
113.85°E) and those with weak OMI SO₂ signal such as 224
Yexian (33.62°N, 113.35°E). Figure 2 shows the SO₂ pro- 225
files observed on 04/15/2008. The destination, Shangcai 226
(33.25°N, 114.26°) had a moderate OMI SO₂ pollution, and 227
relatively high SO₂ concentration (up to 1.5 ppbv) was 228
observed at high altitudes, ~4000 m. Table 1 presents a 229
statistical analysis of ambient SO₂ concentrations averaged 230
in 500 m layers from the surface to 4000 m. Over Chan- 231
gyuan (April 4th and 5th), we observed up to 7 ppbv SO₂ at 232
2000 m, while over Yexian (April 18th and 22nd) the 233
ambient SO₂ concentration was below the detection limit 234
(~0.3 ppbv). The results were consistent with the OMI SO₂ 235
maps. During the descents over the CGO airport, relatively 236
high concentrations of SO₂ were observed consistently, 237
pointing to the urban area of Zhengzhou as a stationary 238
source of the SO₂ pollution. High concentrations of SO₂ 239
were found below ~500 m altitude, implying that the sub- 240
stantial amount of ambient SO₂ was concentrated within 241
the PBL. 242

[15] A summary of flight routes (auxiliary material Figure S1) 243
shows a relatively homogeneous sampling over the prov- 244
ince, with the exception of mountainous northwest region, 245
where complex terrain makes spiral flights unsafe.¹ In April 246
2008, the monthly mean daily average temperature in 247
Zhengzhou was relatively stable at 16.0 ± 5.0 C° (data from 248
www.wunderground.com), so we assumed the sulfur emis- 249
sions from coal burning for electricity generation, domestic 250
heating, and cooking did not change dramatically during the 251
campaign. Therefore, we selected the region covered by the 252
research flights (33.0° to 35.5°N, 112.5° to 115.5°E, here- 253
after named the campaign area) and calculated the campaign 254
average SO₂ profile from the airborne measurements 255
(Figure 3). The integral of SO₂ with respect to altitude gives 256
a mean SO₂ column content of 0.73 Dobson Unit (DU, 257
1 DU = 2.69 × 10¹⁶ molecules/cm²). The majority of SO₂ 258
was found in the PBL and in the lower atmosphere with an 259
effective scale height of SO₂ of ~800 m, from surface up to 260
~1800 m. The mean profile also showed substantial amounts 261
of SO₂ aloft in the FT, where atmospheric SO₂ has a longer 262
lifetime, and is more likely to be converted to sulfate aerosols. 263
The relatively strong winds in the FT transport sulfurous 264
pollutants over greater distances, and SO₂ in the FT has a 265
greater impact on large-scale air quality and climate. 266

¹Auxiliary materials are available in the HTML. doi:10.1029/2011JD016473.

Table 1. Summary of Research Flights^a

Flight ID	Date	Spiral				Descent			
		Altitude (m)	SO ₂ Conc. (ppbv)			Altitude (m)	SO ₂ Conc. (ppbv)		
			Mean	Stdev	Median		Mean	Stdev	Median
t1.6	4/4/2008	1000	4.67	0.29	4.67	500	7.16	0.87	7.02
t1.7		1500	3.74	0.27	3.84	1000	5.14	3.38	3.13
t1.8		2000	2.71	0.34	2.63	1500	1.31	0.33	1.25
t1.9		2500	1.72	0.26	1.65	2000	1.00	0.06	1.01
t1.10		3000	1.79	0.32	1.81				
t1.11		3500	1.23	0.22	1.21				
t1.12		4000	0.87	0.06	0.89				
t1.13									
t1.14	4/5/2008	1000	4.81	0.32	4.93	500	9.48	4.18	8.51
t1.15		1500	6.08	0.75	5.94	1000	2.07	1.42	1.29
t1.16		2000	7.31	0.13	7.34	1500	0.97	0.52	1.24
t1.17		2500	5.96	1.06	6.26	2000	0.20	0.06	0.21
t1.18		3000	2.29	0.83	2.00				
t1.19		3500	1.50	0.07	1.49				
t1.20		4000	1.08	0.16	1.07				
t1.21									
t1.22	4/15/2008	1000	3.00	0.12	2.95	500	4.99	1.86	6.00
t1.23		1500	1.86	0.79	1.58	1000	0.88	0.72	0.52
t1.24		2000	1.07	0.04	1.08	1500	0.51	0.05	0.50
t1.25		2500	0.91	0.07	0.91	2000	0.54	0.03	0.53
t1.26		3000	0.93	0.18	0.92				
t1.27		3500	1.28	0.09	1.24				
t1.28		4000	1.34	0.05	1.35				
t1.29									
t1.30	4/16/2008	1500	0.29	0.12	0.31	500	3.62	1.52	4.35
t1.31		2000	0.18	0.05	0.19	1000	0.42	0.05	0.41
t1.32		2500	0.23	0.03	0.23	1500	0.16	0.23	0.11
t1.33		3000	0.21	0.10	0.22	2000	0.15	0.08	0.19
t1.34		3500	0.32	0.09	0.35				
t1.35									
t1.36	4/18/2008	1500	-0.06	0.06	-0.04	500	3.30	0.35	3.40
t1.37		2000	-0.08	0.04	-0.07	1000	1.38	1.08	1.63
t1.38		2500	-0.11	0.08	-0.10	1500	0.00	0.03	-0.01
t1.39		3000	0.04	0.11	0.09	2000	0.03	0.05	0.02
t1.40		3500	0.25	0.13	0.19				
t1.41									
t1.42	4/20/2008	3000	0.98	0.05	0.98	500	4.14	1.71	4.65
t1.43		3500	1.08	0.17	1.12	1000	0.29	0.15	0.22
t1.44		4000	1.98	0.44	1.79	1500	0.22	0.02	0.21
t1.45		2000	0.70	0.31	0.58				
t1.46									
t1.47	4/22/2008	1500	-0.14	0.02	-0.14	500	3.82	0.10	3.83
t1.48		2000	-0.11	0.07	-0.06	1000	1.04	1.13	0.41
t1.49		2500	-0.14	0.09	-0.19	1500	0.35	0.09	0.38
t1.50		3000	-0.09	0.09	-0.04	2000	0.38	0.07	0.36
t1.51		3500	-0.03	0.04	-0.03				
t1.52		4000	-0.05	0.05	-0.04				

^aNegative values are beyond the detection limit of 43C SO₂ analyzer. Spirals were conducted over Changyuan (April 4 and 5), Shangcai (April 15), Suiping (33.15°N, 113.95°E, April 16), Yexian (April 18 and 22), and Weishi (34.41°N, 114.17°E, April 20). Descents were all conducted over CGO airport.

[16] Figure 3 also shows all the SO₂ measurements from the campaign. The distribution of SO₂ measurements exhibits large variability, especially between 1000 and 3000 m, which is above the typical PBL height during spring in central China. During the campaign, we frequently observed isolated SO₂ plumes in the FT, and the statistics of SO₂ concentrations aloft were greatly influenced by these plumes. All of the research flights were conducted under calm and stable weather conditions without strong convection, so the FT SO₂ plumes were likely related to up wind or large-scale vertical transport. Similar characteristics were observed in Mid-Atlantic region of the U.S. [Hains *et al.*, 2008; Taubman

et al., 2006]. In Table 2, we summarize the FT SO₂ plumes observed during the campaign. To study the transport processes, we calculated 72-h back-trajectories using the NOAA Hybrid Single Particle Lagrangian Integrated Trajectory Model (HYSPLIT, <http://www.arl.noaa.gov/ready/hysplit4.html>), with the Global Data Assimilation System (GDAS) meteorological fields. Time and height of the observed SO₂ plumes were utilized as release time and release height (plume height ± 500 m). For example, Plume 1 demonstrated a stagnant case with air circulating within a radius of 400 km. The deep polluted dry air mass indicated that the lower atmosphere was well mixed. The back-trajectory calculation for plumes 2 and 5 did not demonstrate an upward lifting

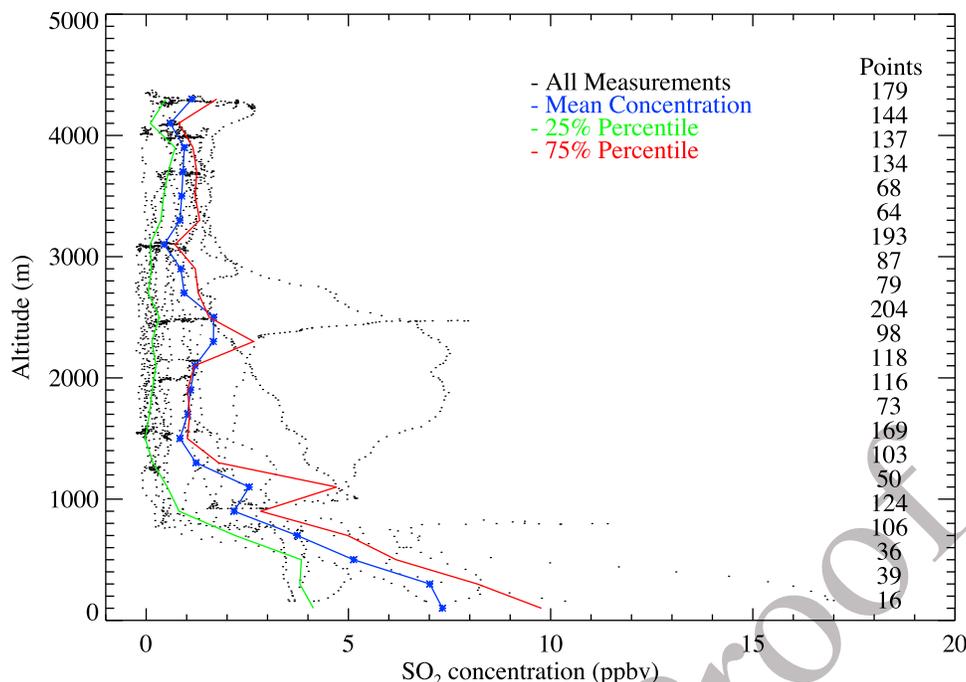


Figure 3. In situ measurements and the mean SO₂ profile. The mean profile is calculated by averaging every 200 m from the surface to 4500 m. Green and red lines indicate the 25th and 75th percentiles. The numbers on the right are the number of valid SO₂ data points within the 200 m layer. The concentration of SO₂ decreases from the surface to ~1800 m with an effective scale height of 800 m, but strata of SO₂ were frequently observed aloft.

292 motion, implying the sub-grid convection could be impor- 312
 293 tant. Figure 4 presents the case study of Plume 3, which 313
 294 existed at a high elevation with high RH value. The back- 314
 295 trajectory suggested a case of long-range transport with what 315
 296 appeared to be isentropic lifting from southern China. The air 316
 297 mass was lifted from the surface to 4200 m and transported 317
 298 around 1500 km in 72 h. Based on GPS data, this plume had a 318
 299 minimum size of 22 km in width and 1000 m in depth. With a 319
 300 mean concentration of 2 ppbv, the total mass of atmospheric 320
 301 SO₂ in the plume was estimated to be at least 1.5 tons. The 321
 302 substantial amount of SO₂ in FT shows the importance of 322
 303 studying large-scale lifting to understand inter-continental 323
 304 transport of pollutants from East Asia. 324

305 **4. Evaluation of OMI SO₂ PBL Products**

306 [17] OMI SO₂ products proved useful for research flight 327
 307 planning during the campaign, and in this section we quanti- 328
 308 tatively evaluate the products. To estimate the accuracies and 329
 309 characterize the limitations of these OMI SO₂ retrievals (both 330
 310 archived operational OMISO2 PBL product and new off-line 331
 311 research ISF product), independent coincident column 332

measurements were needed. The ceiling of in situ measure- 312
 313 ments conducted during the aircraft campaign was ~4500 m, 314
 315 well into the free troposphere. These vertical SO₂ profiles 316
 317 showed that during this campaign, significant amount of 318
 319 ambient SO₂ was found within the PBL (<1 km). Therefore 320
 321 the integration of aircraft vertical profiles can be used to vali- 322
 323 date OMI satellite retrievals. 324

[18] We compared the in situ SO₂ vertical columns with 319
 320 collocated archived OMISO2 PBL values and with the research 321
 322 ISF retrievals using prescribed SO₂ shapes with center of mass 323
 324 altitude (CMAs) similar to the measured in situ profiles over 325
 326 the aircraft spiral locations and their surrounding areas. Fol- 327
 328 lowing the comparison approach described earlier [Krotkov 329
 330 *et al.*, 2008], the average values of the nearest eight pixels in 331
 332 30 km radius of spiral locations were compared with the cor- 333
 334 responding vertically integrated in situ SO₂ columns. The 335
 336 uncertainties of the satellite measurements were estimated as 337
 338 the largest of the OMI average background noise (i.e., 0.62 DU 339
 340 for 8 pixel mean error for the PBL product [Krotkov *et al.*, 341
 342 2008] and standard deviations of the eight nearest pixels). 343
 344 Note that this error estimate approach yields an upper limit, 345

t2.1 **Table 2.** Summary of SO₂ Plumes Observed in the FT

t2.3	Number	Date	Time (UTC)	Location	Conc. (ppbv)	Altitude (m)	Size (km)	RH ^a (%)
t2.4	1	4/15/2008	6:05–6:13	33.85N°, 114.50E°	2.0~8.0	2500~4000	30	30
t2.5	2	4/16/2008	7:00–7:07	34.36N°, 114.31E°	1.3	3700	26	75
t2.6	3	4/20/2008	3:34–3:42	35.26N°, 114.70E°	2.6	4200	22	90
t2.7	4	4/20/2008	4:00–4:13	34.54N°, 115.20E°	2.2	4300	18	100
t2.8	5	4/22/2008	6:53–7:03	33.93N°, 113.11E°	1.6	3700	22	20

t2.9 ^aRH, relative humidity.

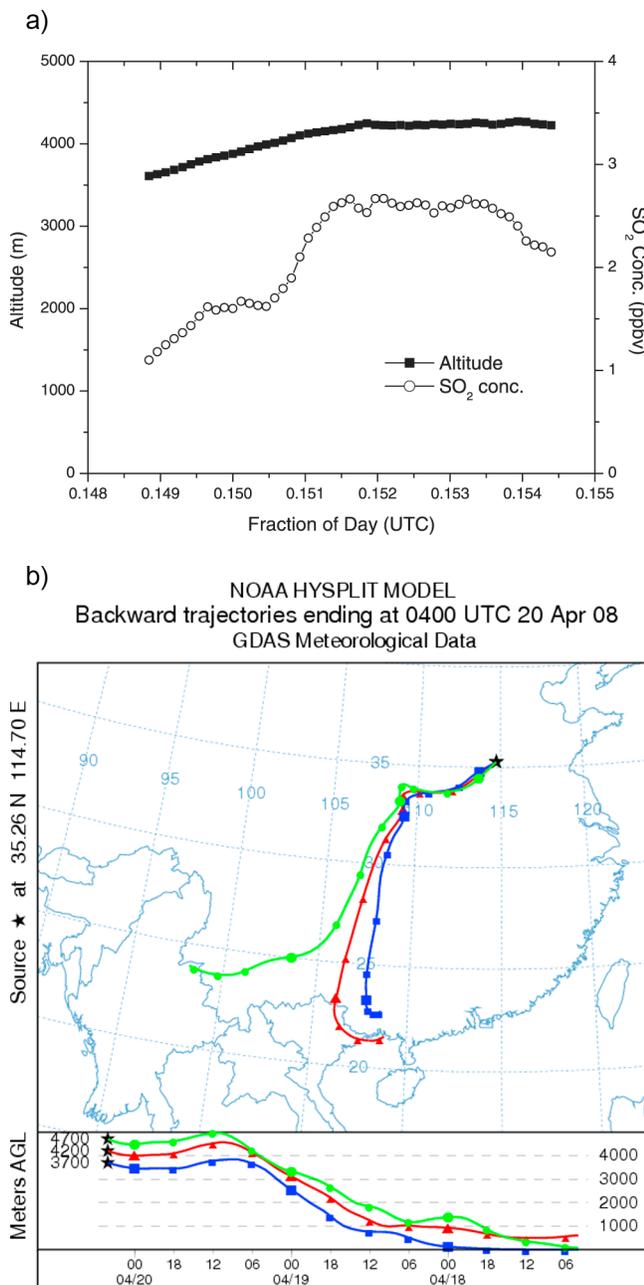


Figure 4. SO₂ plumes in FT and the HYSPLIT back trajectory. (a) Ambient SO₂ concentration versus aircraft altitude; (b) 72-h back trajectory (ending at 0400 UTC, 04/20/2008) of Plume 3. These high altitude plumes affect the OMI retrieval and can be transported relatively long distance.

333 because it includes the natural variability of SO₂ spatial dis-
334 tribution as part of the error.

335 [19] To estimate the error of total column integrated from
336 the in situ concentrations, we need to account for the fact that
337 during the campaign over Henan, many spirals were restricted
338 to altitudes above 1000 m, therefore concentrations below this
339 altitude (down to the ground) were not simultaneously mea-
340 sured. The column amount missed by the aircraft measure-
341 ments below ~1000 m usually accounts for 10–30% of the

total SO₂ PBL column [Lee et al., 2009], and it makes the
largest contribution to the error budget of aircraft measured
total columns. To compensate for these missing columns, we
used the in situ SO₂ measurements obtained during the aircraft
landings at the airport, whose location is different from those
of the spirals. In doing so, we assumed that the partial column
SO₂ amount in the lowest part of the atmosphere is homoge-
neous, and therefore is the same as the mean value observed
during descent on the same day. Under this assumption the
uncertainty of the total vertical SO₂ column is estimated to be
half of the added partial column, as the distance between the
descent and spiral is less than 200 km. A certain degree of
correlation of PBL columns between the two locations is
expected.

[20] We compared SO₂ columns estimated using in situ
measurements and OMISO2 PBL products as the Ordinary
Least Squares (OLS) linear regression (auxiliary material
Figure S2). The OLS slope (0.16) was low, indicating that
OMI greatly underestimated the SO₂ PBL column contents.
The OMISO2 PBL products have variable bias, which need to
be removed empirically [Fioletov et al., 2011]. Here we added
0.4 DU to all OMISO2 PBL data to make them physically
meaningful, and compared the results with in situ measure-
ments (Figure 5a). The comparison showed a strong correla-
tion ($r = 0.84$), higher than that obtained from validation
studies over North America, but the slope (0.54) was still
lower than the previously reported slope found in the compar-
isons of improved OMI products during INTEX-A and
INTEX-B campaign [Lee et al., 2009]. In that study, OMI
PBL data were post-corrected applying local air mass factor
(AMF) calculated using monthly SO₂ profile shapes and
aerosol climatology from the global Goddard Earth Observing
System (GEOS-CHEM) chemical transport model. The correc-
tion resulted in typically reduced SO₂ values over oceans,
including INTEX-A and INTEX-B regions. However, over
China the local AMFs [Lee et al., 2009] were close to the
operational value ~0.4 [Krotkov et al., 2008]. Therefore, no
local AMF correction was applied in this study. The OMI PBL
products underestimated the tropospheric SO₂ column by
~50% likely due to: 1) systematic negative biases in OMISO2
PBL values when the satellite field of views were cloud con-
taminated; 2) reduced satellite measurement sensitivity to SO₂
in the lowest levels due to the presence of aerosols, and 3)
spatial averaging of local SO₂ plumes over large OMI pixel
size. Low visibility (high aerosol concentrations) conditions
were common during our research flights. Visibility observa-
tions (available at www.wunderground.com) at the CGO air-
port at 9 A.M. (local time) from March 28 to April 26 2008
showed most of the flight days with visibility less than 5 km.
The OMI instrument has lower sensitivity to the SO₂ close to
the surface compared to those in the upper atmosphere. The
presence of aerosols above or co-located with the SO₂ layer
would further reduce this sensitivity, leading to an SO₂
underestimate, since aerosol effects are not accounted for in
the operational PBL products. During the EAST-AIRE cam-
paign, the local AMFs could be reduced to ~0.2, half of the
operational AMF (0.4), due to dust aerosols [Krotkov et al.,
2008], implying the OMISO2 PBL products could be dou-
bled to compensate the underestimate of AMFs, while no
correction of AMFs was conducted in this study.

[21] To reduce these random errors, we averaged all OMI
PBL pixel values over the campaign area for April 2008.

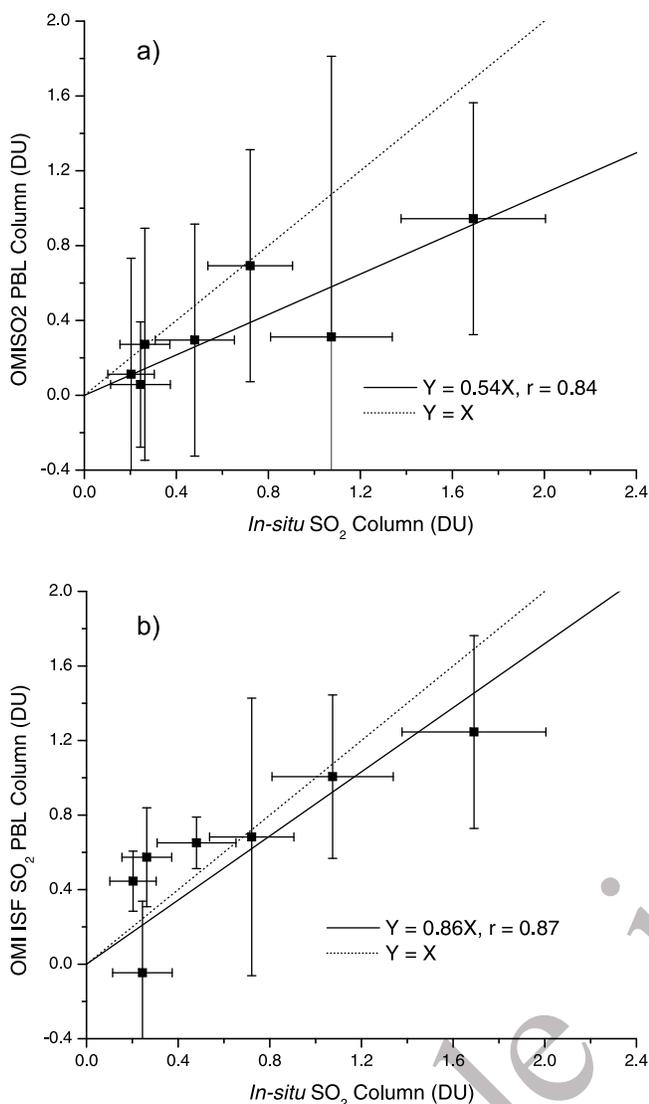


Figure 5. Evaluation of in situ measurements with (a) modified OMISO2 PBL (+0.4 DU) and (b) ISF SO₂ Column Content. The dotted line represents the Y = X line. X and Y error bars describe the uncertainties of aircraft observation and OMI PBL column respectively. Solid line presents the OLS linear regression passing through zero. Both retrievals capture the average column content, but the ISF algorithm shows substantially better fit to individual daily observations.

404 The resulted regional monthly mean ($\pm\sigma$) SO₂ column was
 405 0.63 (± 0.26) DU, close to 0.73 DU obtained from averaging
 406 all aircraft measurements. This suggests that averaged over
 407 a month and a large region, the OMI PBL data can capture
 408 the PBL SO₂ column contents better than them for individual
 409 days.

410 [22] We also compared the integrated in situ data with off-
 411 line ISF retrievals (Figure 5b). The ISF algorithm demon-
 412 strated similar correlation ($r = 0.87$), but less noise and better
 413 evaluation (slope = 0.86) of PBL SO₂ columns, compared
 414 with the operational OMISO2 products. Systematic bias was
 415 not observed, though our comparisons indicate that current
 416 ISF retrievals still have low bias with respect to the airborne

observations over central China. The underestimate could
 be caused by the similar reasons discussed above, so the
 algorithm needs to be improved in this respect possibly
 through incorporation of tropospheric aerosols information
 in future.

5. Numerical Simulations of SO₂ Over Central China

5.1. WRF-CMAQ Model Setup

[23] In this study, we used the WRF V3.1 model [National
 Center for Atmospheric Research, 2010] to generate meteo-
 rological fields for the CMAQ model. Figure 6 shows the two
 WRF domains in a Lambert projection. The coarse domain
 (30 km grid cells) covered the central and eastern part of
 China, where most of the population and industry were
 located, and the nested domain (10 km grid cells) focused on
 the campaign region. The U.S. Geological Survey (USGS)
 24-category data were used to determine the terrain and land
 use. We used the NCEP Final Operational Model Global
 Tropospheric Analyses (FNL) (<http://dss.ucar.edu/datasets/ds083.2>)
 as initial and boundary conditions. The NCEP FNL
 are on $1^\circ \times 1^\circ$ grids with 26 vertical levels from 1000 to
 10 hPa with a time frame of 6 h. The major physics options
 used in the WRF simulation included Thompson microphysics
 scheme [Thompson *et al.*, 2006], YSU boundary layer scheme
 [Hong and Lim, 2006], Kain-Fritsch (new Eta) cumulus
 scheme [Kain, 2004], Monin-Obukhov surface-layer scheme
 [Foken, 2006] and Noah land-surface scheme [Ek *et al.*, 2003].
 The model was run with 35 vertical layers from the surface to
 50 hPa with the first 12 layers in the PBL, and re-initialized
 every 5 d to reduce simulation errors. MCIP V3.5 was applied
 to process the WRF outputs to create CMAQ-ready meteo-
 rology inputs [Byun and Ching, 1999].

[24] We used the CMAQ version 4.6 (released in September
 2006) [Byun and Schere, 2006] to conduct a 45-d simulation

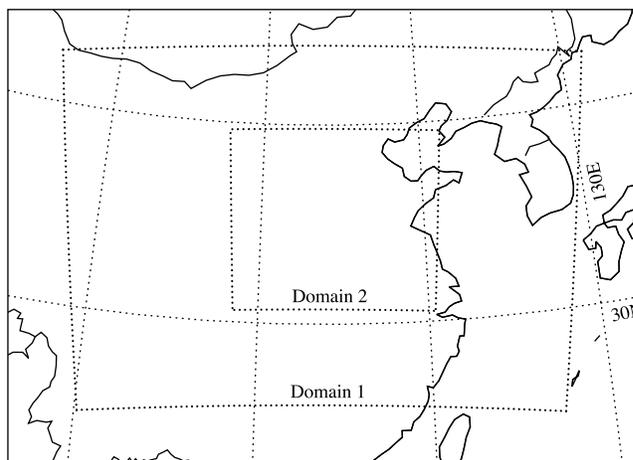


Figure 6. Domains of the WRF simulation. Both the coarse and nested domains are centered at Zhengzhou City (34.8°N , 113.7°E). The coarse domain has 97×69 grids and 30 km resolution, and the nested domain has 117×102 grids and 10 km resolution. The CMAQ domains are 2 grid cells smaller on each side of the WRF domains to eliminate the lateral effect.

451 from 03/13/2008 to 04/26/2008, with the first 15 days as
452 spin-up. CMAQ was run with a coarse and nested domain
453 with the same horizontal and vertical resolutions as the WRF
454 simulation. Chemical initial and boundary conditions for the
455 coarse domain were obtained from the Regional Acid
456 Deposition Model, version 2 (RADM2) concentration pro-
457 files [Community Modeling and Analysis System, 2007;
458 Stockwell et al., 1990]. The SAPRC99 scheme and the 4th
459 generation aerosols scheme (AE4) were selected as the gas-
460 chemistry mechanism and aerosol modules respectively. The
461 Regional Acid Deposition Model (RADM) based cloud
462 processor with the asymmetric convection model (ACM) was
463 applied for the aqueous/cloud chemical mechanism. The
464 CMAQ output files included hourly 3-D fields of gaseous
465 and aerosol species concentrations.

466 5.2. The Emission Inventory

467 [25] We selected the NASA INTEX-B emission inventory
468 (available at <http://mic.greensource.cn/intex-b2006>) for the
469 CMAQ simulations, although there is evidence that emis-
470 sions decreased between 2006 and 2008 [Witte et al., 2009].
471 This latest emission inventory of East Asia was based on the
472 year 2006 with estimates of all major anthropogenic sources
473 [Zhang et al., 2009], including major pollutants (SO_2 , NO_x ,
474 CO , PM_{10} , $\text{PM}_{2.5}$, BC , and OC) and 30 lumped VOC spe-
475 cies for SAPRC-99 chemical mechanism with a resolution
476 of $0.5^\circ \times 0.5^\circ$. Compared with EPA's National Emission
477 Inventory (NEI) database, INTEX-B only contained area
478 emission sources, and lacked point sources, mobiles sources
479 and other geographic information. Therefore, we could not
480 use the Sparse Matrix Operator Kernel Emission (SMOKE)
481 model to create the 3-D emission input data for CMAQ, and
482 had to create our own emission input data through the fol-
483 lowing steps. First, we incorporated the 2008 NH_3 emission
484 prediction from the REAS program (available at <http://www.jamstec.go.jp/frsgc/research/d4/emission.htm>) as supplement.
485 Second, INTEX-B had each pollutant calculated individually
486 for four sectors: Electricity Generation, Industry, Residential
487 Emission and Transportation, and we allocated them into two
488 groups: the Electricity Generation and Others. Since stack
489 parameters such as plume exit velocity and plume exit tem-
490 perature were not available, Electricity Generation emissions
491 were located 200 m above the surface as an approximation
492 of average stack height and plume rise. Third, we speciated
493 the emission data of NO_x and $\text{PM}_{2.5}$ into NO_2 , NO , sulfate
494 and nitrate aerosols to accommodate the SAPRC99 mecha-
495 nism. Last, we averaged the yearly emission values into
496 hourly values arithmetically, and allocated them into the
497 CMAQ grid cells through bilinear interpolation, to create a
498 3-D emission input data with constant values for all chemical
499 species.

501 [26] The INTEX-B inventory estimated total emission of
502 SO_2 for China was 31.0 Tg per year with $\pm 12\%$ reported
503 uncertainty, and the arithmetic mean daily SO_2 emission
504 was 84.9 kT, which was close to 84.8 kT and 78.9 kT from
505 the arithmetic mean from March and April emissions
506 respectively [Zhang et al., 2009]. These values implied that
507 the seasonality of SO_2 emissions was negligible during
508 spring in China. We ignored the diurnal cycle of emissions
509 by assuming constant emission rates since there was no
510 available information on diurnal variation. For demonstra-
511 tion, we presented the resulting SO_2 emission maps

(auxiliary material Figure S3), which correspond well to 512
the location of cities, populations and industrial centers in 513
China such as YRD. This spatial accuracy was crucial to the 514
CMAQ simulations. The manipulation of INTEX-B emis- 515
sion inventory created CMAQ emission input files without 516
seasonal and diurnal variations. The effects of using 2006 517
INTEX – B emissions on CMAQ simulations for 2008 will 518
be discussed in section 7. 519

520 5.3. Modification of the CMAQ System

[27] Version 3.5 of MCIP was used to ingest the WRF 521
output and create meteorological input files for the CMAQ 522
model. We modified MCIP to write out the percentage of 523
each WRF grid cell that is urban, and the new urban fraction 524
variable was used to calculate vertical diffusion in CMAQ 525
[Castellanos, 2009]. The SO_2 dry deposition velocity was 526
calculated in the MCIP program. In the default MCIP set- 527
ting, the mean SO_2 dry deposition velocity ($\pm\sigma$) was $0.58 \pm$ 528
 0.07 cm/s over the campaign area, which is substantially 529
higher than $0.2 \sim 0.4$ cm/s measured over northern China 530
[Sorimachi et al., 2003; Sorimachi and Sakamoto, 2007] and 531
other areas [Clarke et al., 1997]. To decrease the simulated 532
dry deposition rate, we set the mesophyll resistance of SO_2 533
from 0 (default value) to 8000 s/m [Pfanzen et al., 1987] in the 534
MCIP model. Similar modification successfully decreased 535
the CO dry deposition velocity from ~ 0.4 cm/s to ~ 0.1 cm/s 536
[Castellanos et al., 2011]. 537

[28] The CMAQ model was modified to facilitate our 538
analysis of the sulfur budget. To estimate the export of S 539
(sulfur) from China, the CMAQ code was modified to output 540
the flux of each species due to horizontal advection entering 541
and leaving the domain through changing the advection 542
scheme from the Yamartino (HYAMO) scheme to the new 543
Piecewise Parabolic Method (HPPM) scheme [Loughner, 544
2011; Loughner et al., 2011]. This allowed pollutant fluxes, 545
deposition, and emissions to be studied collectively, and we 546
applied these results to study the S budget and export in 547
section 6.3. To investigate influences on SO_2 chemistry from 548
these modifications, we conducted three sensitivity runs: (1) 549
default mesophyll resistance (MR), and HYAMO advection 550
scheme; (2) default MR and HPPM scheme; and (3) updated 551
MR and HPPM scheme. They are named NoMR_HYAMO, 552
NoMR_HPPM, and MR_HPPM respectively. 553

[29] To isolate effects of modifying SO_2 mesophyll resis- 554
tance, the difference between results from NoMR_HPPM 555
and MR_HPPM was calculated. We selected data over 556
the campaign area (mostly grassy plain with good vegetation 557
coverage), on days with calm weather and no precipitation to 558
minimize the effects of SO_2 wet deposition. The monthly 559
average dry deposition velocity ($\pm\sigma$) of the MR_HPPM run 560
was 0.24 ± 0.06 cm/s, $\sim 60\%$ lower than the NoMR_HPPM 561
run with the value of 0.58 ± 0.07 cm/s. We also investigated 562
the SO_2 dry deposition flux, calculated in the CMAQ model. 563
The monthly mean value ($\pm\sigma$) decreased from $0.093 \pm$ 564
 0.039 $\mu\text{g}/\text{m}^2$ s with NoMR_HPPM to 0.045 ± 0.031 $\mu\text{g}/\text{m}^2$ s 565
with MR_HPPM, a 52% reduction. The MR_HPPM value 566
is around 50% less than a lab experiment conducted in 567
northern China, but the SO_2 dry deposition process is influ- 568
enced by other factors such as soil type and meteorological 569
conditions [Sorimachi and Sakamoto, 2007]. These results 570
show that a better parameterization of SO_2 dry deposition 571
is needed for future development of the CMAQ model. 572

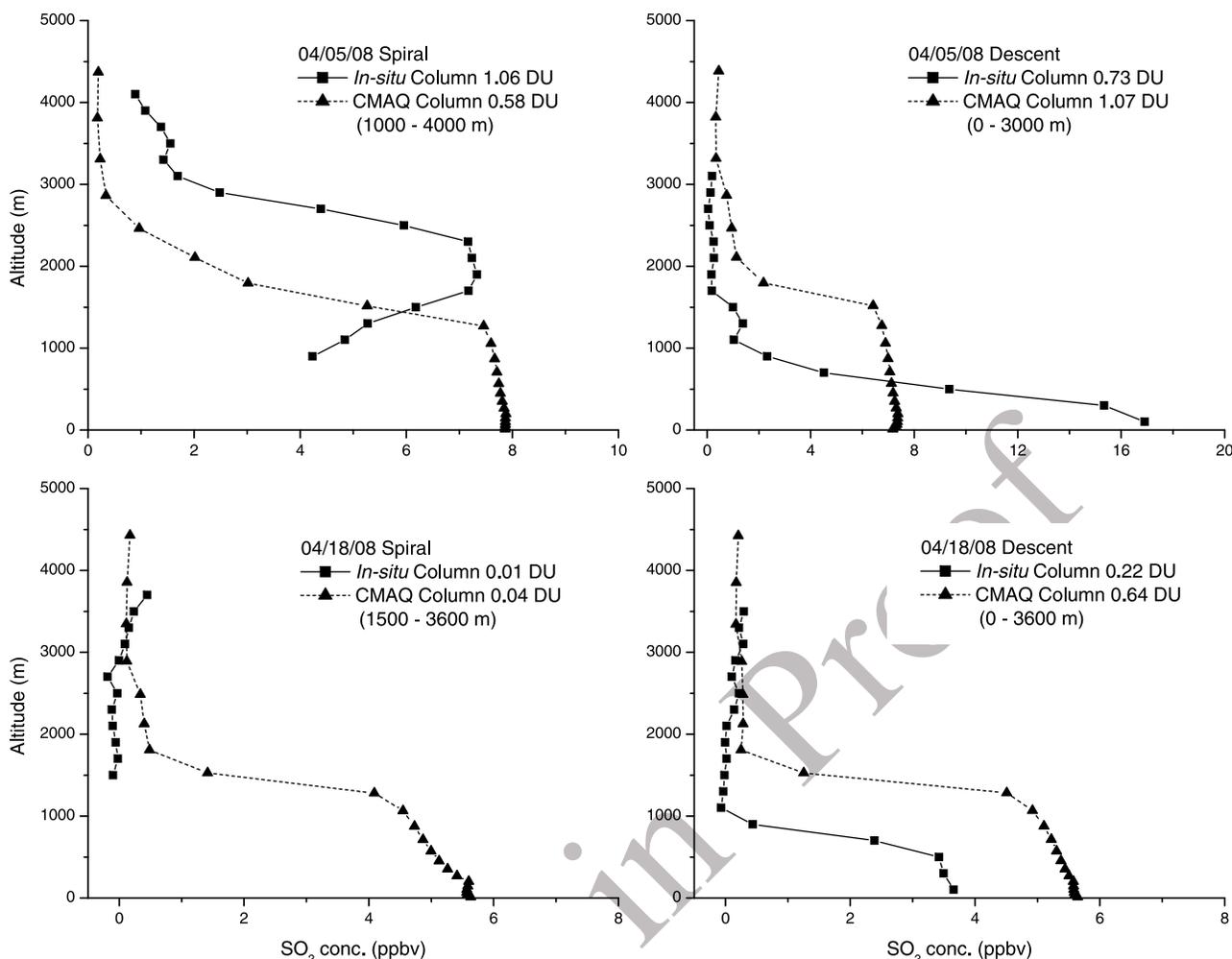


Figure 7. Comparison of aircraft measurements and CMAQ simulations. The in situ measurements are obtained by averaging the aircraft SO_2 profiles every 200 m from surface to 4500 m. The CMAQ SO_2 column contents are calculated by integrating the SO_2 at the same altitudes as aircraft measurements. The model cannot fully resolve the subgrid scale processes responsible for SO_2 plumes aloft.

573 Hereafter in the following sections, we use results from
574 MR_HPPM run and discuss the details of the three sensitivity
575 runs in section 6.1.

576 6. Results of CMAQ Simulation and Discussion

577 6.1. Evaluation of CMAQ Simulations

578 [30] We compared two days' flight data with the CMAQ
579 results, and found that the SO_2 profiles were not precisely
580 simulated especially within the PBL (Figure 7). We also
581 checked CMAQ simulations for the FT SO_2 plumes dis-
582 cussed in Table 2, and the model did not reproduce these
583 plumes well. This implies that CMAQ (with 10 km resolu-
584 tion) had difficulty reproducing the vertical SO_2 altitude
585 profiles at specific times and locations. Other studies have
586 shown that CMAQ has similar problems with vertical pro-
587 files of trace gases over the eastern U.S. [Castellanos *et al.*,
588 2011; Lee *et al.*, 2011]. These issues were probably due to
589 sub-grid scale convective lifting, and the model resolution of

10 km was apparently inadequate to simulate the vertical
590 mixing [Loughner *et al.*, 2011].

591 [31] Other sources of uncertainty include the precision
592 of winds. Due to limited resource, we did not use the Four-
593 Dimensional Data Assimilation (FDDA) in WRF simula-
594 tions. Without FDDA, wind errors were expected resulting
595 in uncertainty of CMAQ simulations [Otte, 2008]. SO_2
596 emissions from power plants were estimated at 200 m AGL,
597 however these plumes could rise to several hundred meters
598 under certain weather conditions. Therefore, we focused on
599 comparisons of SO_2 column contents of each research flight
600 to the corresponding SO_2 columns from CMAQ simulations
601 in Figure 8. The CMAQ simulations had a moderately strong
602 correlation ($r = 0.62$) with the integrated aircraft measure-
603 ments, with a slope of 1.33. This indicated that CMAQ pro-
604 vided a reasonable representation of the spatial/temporal
605 variation of total SO_2 loading over the campaign area and
606 period, and $\sim 30\%$ overestimate was observed.
607

608 [32] Figure 9 compares CMAQ monthly mean SO_2 profile
609 over the campaign region, with the mean aircraft campaign
609

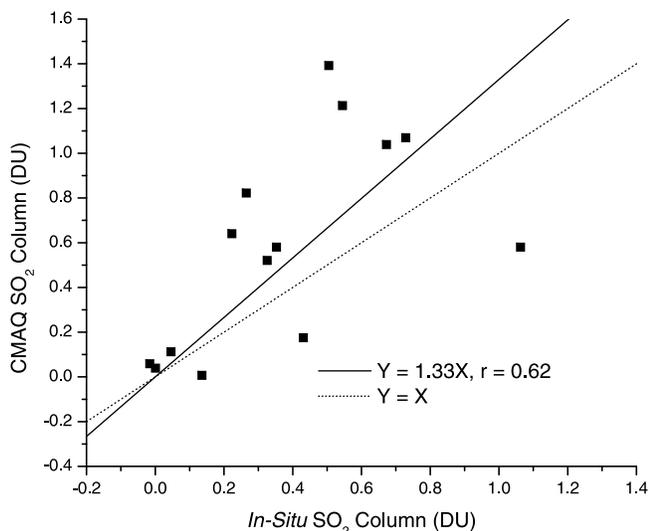


Figure 8. Evaluation of in situ and CMAQ SO₂ column contents. The CMAQ SO₂ column was calculated over corresponding altitude of research flights. The dotted line represents the $Y = X$ line. The solid line presents the OSL linear regression, and there were 7 flights and 14 profiles giving 6–13 degrees of freedom.

610 profile. The monthly mean MR_HPPM profile underestimated
611 the ambient SO₂ near the surface, overestimated the SO₂
612 between 600 and 2200 m, and underestimated the SO₂ above
613 2500 m. The total SO₂ column was ~38% higher than the value
614 obtained during the aircraft campaign. All the three sensitivity
615 runs over predicted the SO₂ columns. Based on differences of
616 the NoMR_HPPM and MR_HPPM cases, it was confirmed that
617 decrease of the SO₂ dry deposition increased the ambient SO₂
618 column content. Due to little SO₂ emission in the upwind
619 western China, the comparison between the NoMR_HYAMO

and NoMR_HPPM cases implied that the HPPM advection
620 scheme transported less SO₂ out of the domain. The default
621 CMAQ set up (NoMR_HYAMO case) demonstrated the best
622 estimate of total SO₂ column content, while the modified case
623 (MR_HPPM) had the best estimate of SO₂ concentration near
624 the surface. Besides uncertainties in meteorological and
625 emission data, the underestimate/overestimate within/above
626 the PBL implied that the CMAQ model might mix the lower
627 atmosphere too fast to transport the pollutants out of the PBL,
628 also observed in the eastern U.S. [Castellanos *et al.*, 2011].
629

[33] To summarize, the CMAQ simulations demonstrated
630 reasonable accuracy (~30% overestimate) in reproducing
631 the SO₂ column content observed during the research flights,
632 but the SO₂ model profiles had substantial differences with
633 respect to the measured SO₂ vertical distributions. More
634 accurate model representation of pollutant vertical distribu-
635 tion will require better parameterization of vertical transport
636 and mixing in future. The comparison of monthly mean SO₂
637 profiles from sensitivity runs showed that the CMAQ model
638 overestimated the SO₂ column content by 25 ~ 40% from
639 difference sensitivity runs. The discrepancies are probably
640 caused by the decrease from the 2006 emission inventory
641 used in this CMAQ simulation. The recent decreasing trend
642 in Chinese SO₂ pollution has been observed in number of
643 studies [Li *et al.*, 2010b; Lu *et al.*, 2010; Witte *et al.*, 2009]
644 and attributed to wide installation of Flue Gas Desulfuriza-
645 tion (FGD) equipment on coal burning power plants.
646

6.2. SO₂ Chemistry and Lifetime

[34] To investigate the conversion of SO₂ to other sulfur-
648 compounds such as sulfate aerosols, we estimated lifetime of
649 tropospheric SO₂. A simple box model was applied to the
650 nested domain (area with $\sim 1.25 \times 10^6$ km²), and the SO₂
651 lifetime was calculated as $\tau_{SO_2} = \frac{\text{Loading}}{\text{Emission}}$. The average SO₂
652 emission was 4.3×10^{-3} mol/km² s based on the INTEX-B
653 emission inventory. The campaign mean SO₂ loading from
654

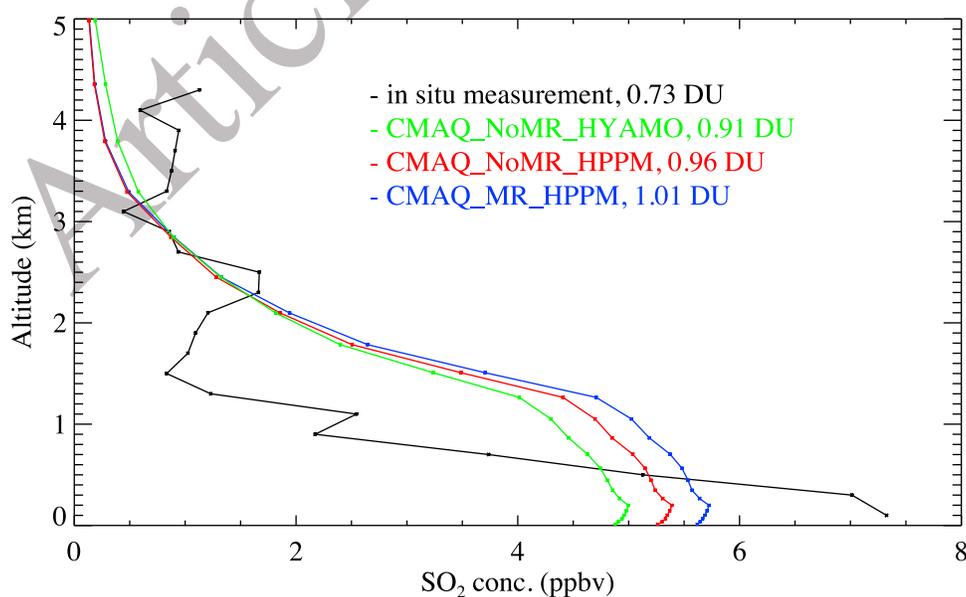


Figure 9. Comparison of monthly mean profiles of in situ measurements and CMAQ sensitivity runs. The SO₂ vertical column amount is the integral from the surface to 5000 m.

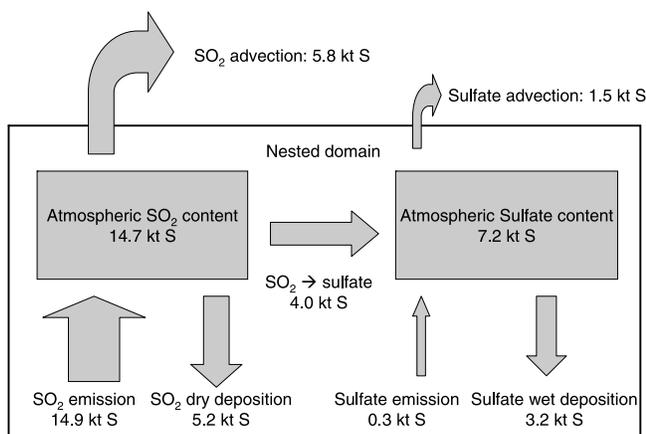


Figure 10. Budget of sulfur-compounds over the central China. The units for reservoirs are kT = (1 kT = 10^6 kg), or kT/d for fluxes. All the data are the average daily value from the month-long simulations.

655 the aircraft campaign was 326 mol/km^2 , and the SO_2 lifetime
 656 with respect to all losses from the domain (deposition,
 657 chemical transformation, and advection) was 21.0 h. The
 658 lifetime was shorter than the result from a global model
 659 simulation [Lee *et al.*, 2011]. In summer over the Mid-
 660 Atlantic, the observed lifetime for chemical removal and
 661 deposition was $19 \pm 7 \text{ h}$ [Hains, 2007]. April in China is
 662 cooler and drier than summer over the Mid-Atlantic resulting
 663 in low photochemical reactivity, and a longer SO_2 lifetime is
 664 expected. The shortcoming of this approach was that the
 665 ambient SO_2 was assumed being emitted within the box
 666 model region. The size of nested domain was only $\sim 1000 \times$
 667 1000 km , so an air mass with the typical wind speed of 10 m/
 668 s would move through it within one day. Therefore, sub-
 669 stantial transport of sulfur-compounds in and/or out of the
 670 domain was expected.

671 [35] Modification of the CMAQ advection scheme pro-
 672 vided the ability to calculate the advection of pollutants
 673 through the boundaries; therefore we can investigate the
 674 budget of sulfur-compounds (SO_2 and sulfate aerosols). To
 675 simplify the computation, we described all sulfur-com-
 676 pounds in units of kT S. The reacted SO_2 amount, calculated
 677 as $\text{SO}_2(\text{reactive}) = \text{Emission} - \text{Deposition} - \text{Advection}$,
 678 assumed that the SO_2 oxidation product was 100% sulfate
 679 aerosols, i.e., 1 mol SO_2 is equivalent to 1 mol of sulfate.
 680 Figure 10 presents the monthly mean budget of sulfur-
 681 compounds of the nested domain, in units of kT/day. The
 682 daily SO_2 emission was 14.9 kT S and the monthly average
 683 SO_2 content in the domain was 14.7 kT S. The major sinks
 684 of SO_2 include daily dry deposition of 5.2 kT S and advec-
 685 tion of 5.8 kT S out of the domain. The average rate of oxi-
 686 dation of SO_2 was assumed as 4.0 kT S/d, which generated
 687 the same amount of S in sulfate aerosols. The direct sulfate
 688 input from the emission inventory was small, 0.3 kT S. The
 689 discrepancy of sulfate budget (*sources* – *sinks*) is 0.5 kT S,
 690 accounting for less than 3% of the total daily S emission
 691 (14.9 kT S), which could be caused by: 1) assuming the
 692 100% conversion from SO_2 to sulfate aerosols and ignoring
 693 the pathways to form other sulfur-compounds such as sul-
 694 furic acid vapor (H_2SO_4); and 2) neglecting minor

deposition processes such as SO_2 wet deposition and sulfate
 dry deposition. We re-calculated the lifetime of SO_2 as

$$\tau_{\text{SO}_2} = \frac{\text{Loading}}{\text{Emission} - \text{Export}}$$
 Based on Figure 10, the total amount of
 exported SO_2 was 5.8 kT/d, which accounts for $\sim 39\%$ of the
 total emitted SO_2 . Including error analyses, the CMAQ
 simulated SO_2 budget and lifetime was computed for each
 day of the campaign, and the average lifetime ($\pm\sigma$) was
 $38 \pm 7 \text{ h}$. Due to significant advection, the oxidation rate
 was substantially decreased and the SO_2 lifetime is consis-
 tent with results from global model simulations [Lee *et al.*,
 2011], and the case study in northern China with a SO_2
 lifetime of $\sim 2 \text{ d}$ [Li *et al.*, 2010a].

6.3. Estimated Sulfur Transport and Error Analysis

[36] We further investigated the sinks and transport of
 sulfur-compounds. The daily total deposition (wet + dry)
 was 8.4 kT S, accounting for 55% of the total sulfur emis-
 sion. This demonstrated the importance of controlling sulfur
 emissions for mitigating the soil and water acidification in
 central China [Larssen *et al.*, 2006]. Daily, 7.3 kT S were
 transported out of the domain, and the monthly mean ($\pm\sigma$) S
 of daily export was $48 \pm 7\%$ of the total sulfur emission.

[37] Our estimate of atmospheric sulfur export from central
 China is subject to random and systematic errors. The ran-
 dom uncertainty can be estimated from the standard deviation
 of daily sulfur export, 7%. A systematic bias is incurred if
 the month was atypical with respect to pollutant exports.
 April 2008 was observed as a generally wet month, with
 more precipitation than normal, 54 mm in April 2008 com-
 pared with 43 mm for long-term monthly average precipita-
 tion (data from www.wunderground.com). The WRF-
 CMAQ model also reproduced the wet month, therefore our
 approach using CMAQ simulations would overestimate the
 average sulfur wet deposition and underestimate export. We
 assumed conservatively the overestimation as half of the
 mean sulfate wet deposition: 1.6 kT S/d, or 10.5% of total
 S emissions (14.9 kT/d). Finally, errors in the simulated
 vertical distribution of SO_2 (Figure 9) affected the horizontal
 flux. We calculated the difference of average SO_2 concen-
 tration between in situ measurements and CMAQ as
 $+0.5 \text{ ppbv}$, -2.0 ppbv and $+0.5 \text{ ppbv}$ for layers 3000 ~ 4500 m,
 500 ~ 2000 m, and surface to 500 m respectively. The wind
 speeds for these layers were estimated at 15 m/s, 8 m/s and
 5 m/s respectably, based on statistics of zonal mean flow. The
 SO_2 export was proportional to the product of SO_2 concen-
 tration and wind speed. So the CMAQ model had biases of
 $+1.0 \text{ kT S}$, -3.0 kT S and $+0.3 \text{ kT S}$ for these three layers,
 which resulted in a net overestimate of 1.7 kT S in export,
 i.e., +11% of the total sulfur-compound emission. Adding
 these three uncertainties ($\pm 7\%$ and $\pm 11\%$) in quadrature
 provided an estimate of the total uncertainty for our estimate of
 S export. We concluded that 48% ($\pm 13\%$) of the S emitted into
 the atmosphere over central China was exported. The uncer-
 tainties were not Gaussian distributions, but we estimated
 the confidence interval as 90% for the range of 35 ~ 61%.

[38] The nested domain is located in central China, and there
 are limited sulfur emissions in the upwind region of less
 developed western China. The study of pollutant transport
 demonstrated that most of the exported sulfur-compounds
 were transported through the northern and eastern boundary to
 the western Pacific coast (auxiliary material Figures S4a–S4d).

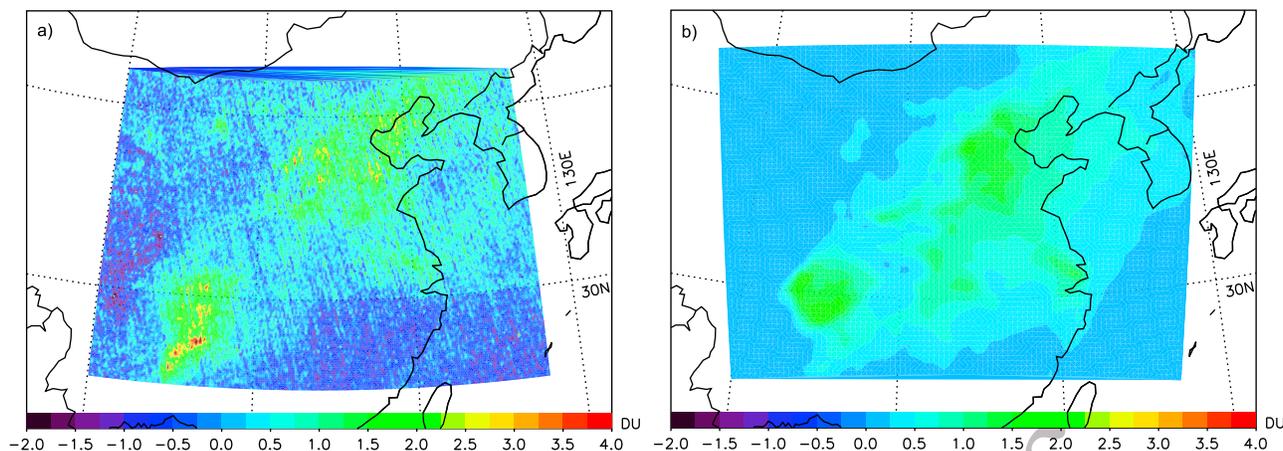


Figure 11. Monthly average SO₂ column maps of the CMAQ simulations and OMI PBL products. (a) OMI SO₂ PBL column, 42 kT in 4.5×10^6 km²; white color describe the location of clouds; (b) CMAQ SO₂ column map, 54 kT in 4.6×10^6 km².

755 Prior studies on S budgets through airborne measurements
756 [Koike *et al.*, 2003] and model simulations [Tan *et al.*, 2002]
757 also reported that around half of the sulfur-compounds were
758 exported to the ocean. Quantifying long-range transport is
759 crucial for understanding the regional/global air quality and
760 climate change due to emissions from East Asia.

761 6.4. Estimate of SO₂ Emissions in Central 762 and Eastern China

763 [39] In section 3, we calculated the monthly mean OMI
764 SO₂ column content over the campaign region as $0.63 \pm$
765 0.26 DU, which agrees with the observed aircraft average
766 SO₂ column. CMAQ overestimated the monthly mean SO₂
767 column over the campaign region with a column content of
768 1.01 DU. The CMAQ high bias could be due to SO₂ being
769 transported from nearby regions into the small campaign
770 region. Because of small size, the OMI instrument had only
771 ~ 150 pixels per day within the campaign region. To reduce
772 the noise and eliminate the influence of SO₂ transport, we
773 selected CMAQ results of coarse domain (Figure 6) to study
774 the total SO₂ burden over central and eastern China. The
775 coarse domain ($\sim 4.6 \times 10^6$ km²) covered most of the indus-
776 trialized regions in China, and emitted ~ 25 Tg SO₂/year,
777 $\sim 80\%$ of the total SO₂ emission in China [Zhang *et al.*, 2009].
778 We sampled CMAQ and OMISO2 PBL columns over the
779 region with OMI reactive cloud fraction less than 0.3 (here-
780 after named cloud-free condition), in order to reduce cloud
781 effects on the satellite retrievals.

782 [40] A daily comparison was conducted on the CMAQ SO₂
783 and OMISO2 columns for 04/05/2008 (auxiliary material
784 Figures S5a and S5b). The CMAQ simulations captured the
785 large-scale SO₂ plumes in the northeast and southeast of
786 China, but missed the plume in the northwest. We found the
787 daily OMI SO₂ map had large variability due to both uneven
788 coverage (sparse data on the east and west side of the domain)
789 and instrument noise (negative value up to -2.0 DU). The
790 CMAQ model overestimated the SO₂ burden by 10% com-
791 pared with the OMISO2 product for 04/05/2008. To reduce
792 the uncertainties of daily data, we presented the monthly
793 average SO₂ column map for both data sets (Figure 11). The

CMAQ model captured the hot spots of SO₂ well over land, 794
but missed the plumes off the coast. Column differences, 795
(OMI – CMAQ) were calculated resulting in a mean differ- 796
ence of -0.16 DU, and the histogram and probability density 797
was analyzed (auxiliary material Figure S6) indicating a 798
negative bias. So the CMAQ simulations systematically 799
overestimated the SO₂ column by ~ 0.16 DU compared with 800
OMISO2 PBL data. The monthly mean SO₂ loadings ($\pm\sigma$) 801
were 54 ± 22 kT and 42 ± 26 kT for CMAQ simulations and 802
OMISO2 PBL products respectively. Because both the 803
OMISO2 products and CMAQ simulations were sampled 804
under cloud-free conditions using OMI radiative cloud frac- 805
tion, these absolute low bias was expected to daily OMISO₂ 806
and CMAQ SO₂ loadings. The monthly mean cloud-free 807
fraction ($\pm\sigma$) was $56 \pm 15\%$. The same approach of sulfur 808
export error analysis was adopted here, i.e., underestimate as 809
half of SO₂ loading filtered out. The systematic low bias was 810
estimated as $+21$ kT and $+16$ kT for CMAQ simulations and 811
OMISO2 PBL products respectively, therefore the mean SO₂ 812
loadings with all uncertainties were 54 ($55 \sim 97$) kT and 813
42 ($37 \sim 89$) kT. The result revealed that CMAQ simulations 814
overestimated OMI SO₂ column contents by $\sim 30\%$, consis- 815
tent with the relative change in SO₂ emissions in China 816
recently [Lu *et al.*, 2010]. 817

[41] As a summary, CMAQ simulations with the INTEX-B 818
emission inventory demonstrated reasonable performance in 819
describing the intensity and spatial distribution of sulfur- 820
compound emissions in China. Both statistical analyses and 821
total SO₂ loading computation indicated that CMAQ over- 822
estimated the SO₂ column contents compared with OMI 823
products, suggesting this might be due to the use of 2006 824
emission inventory; our results are consistent with reported 825
decreasing trends of sulfur emission from 2006 to 2008 in 826
China [Lu *et al.*, 2010; Witte *et al.*, 2009]. 827

7. Discussion and Conclusions 828

[42] Using a combination of in situ and remotely sensed 829
measurement along with numerical simulation, we were able 830
to evaluate the SO₂ concentration and chemistry over central 831
China – one of the most densely populated regions in the 832

833 world. The CMAQ simulations served as a powerful tool to
 834 investigate tropospheric sulfur pollutants. Among the most
 835 important sources of uncertainties in the numerical simula-
 836 tions was the emission inventory, which was developed for
 837 the year 2006. Our campaign was conducted in 2008, and
 838 from 2006 to 2008, especially in order to improve the air
 839 quality for the 2008 Beijing Olympics, sulfur emissions in
 840 China were decreased through stricter regulations on the
 841 usage of high sulfur coal and installations of FGD equipment
 842 in power plants. The decreasing trend has been observed
 843 [Okuda et al., 2011; Witte et al., 2009], so the INTEX-B
 844 emission inventory probably have high bias, resulting in
 845 overestimated anthropogenic SO₂ emissions in the CMAQ
 846 simulations for spring 2008. A comparison with campaign
 847 SO₂ column contents demonstrated that CMAQ overestimated
 848 the tropospheric SO₂ columns over Henan by ~30%, and a
 849 similar overprediction, also ~30%, was obtained through
 850 comparing monthly CMAQ and OMI SO₂ column contents
 851 over central and eastern China. These results are consistent
 852 with the reduction of sulfur emissions observed in China
 853 between years 2006 and 2008.

854 [43] During the campaign, the aircraft instrument fre-
 855 quently observed SO₂ plumes in the FT, important for sat-
 856 ellite retrieval and long-range transport. We found the SO₂
 857 concentrations had high temporal and spatial variability
 858 during spring in China. Comparisons of in situ measure-
 859 ments, the operational OMISO2 PBL products revealed good
 860 agreement, and the new ISF algorithm demonstrated better
 861 performance and will be employed operationally in the
 862 future. The CMAQ model didn't capture the SO₂ vertical
 863 distribution well, probably due to inadequate model resolu-
 864 tion for computing vertical mixing. But the CMAQ simula-
 865 tions agreed with the mean aircraft measurements, with a
 866 ~30% overestimation. With a modified CMAQ advection
 867 scheme, we investigated the budget and transport of SO₂ and
 868 sulfate. The lifetime of SO₂ with respect to all reactions and
 869 removal from the atmosphere was 38 ± 7 h, relatively long
 870 for spring. Due to the slow removal and strong winds in
 871 spring, ~50% (35 ~ 61%) of the total S emitted into the
 872 atmosphere in central and eastern China was transported out
 873 of the domain. Further research in East Asia will help to
 874 improve our knowledge on the effects on regional air quality
 875 and large-scale climate downstream.

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880 References

881 Albrecht, B. A. (1989), Aerosols, cloud microphysics, and fractional cloudi-
 882 ness, *Science*, 245(4923), 1227–1230, doi:10.1126/science.245.4923.1227.
 883 Berglen, T. F., T. K. Berntsen, I. S. A. Isaksen, and J. K. Sundet (2004),
 884 A global model of the coupled sulfur/oxidant chemistry in the tropo-
 885 sphere: The sulfur cycle, *J. Geophys. Res.*, 109, D19310, doi:10.1029/
 886 2003JD003948.
 887 Byun, D. W., and J. E. Ching (1999), Meteorology-Chemistry Interface Pro-
 888 cessor (MCIP) for Models-3 Community Multiscale Air Quality (CMAQ)
 889 modeling system, in *Science Algorithm of the EPA Models-3 Community*
 890 *Multi-scale Air Quality (CMAQ) Modeling System*, EPA/600/R-99/030,
 891 pp. 12.1–12.87, U.S. EPA, Research Triangle Park, N. C.
 892 Byun, D., and K. L. Schere (2006), Review of the governing equations,
 893 computational algorithms, and other components of the models-3 Com-
 894 munity Multiscale Air Quality (CMAQ) modeling system, *Appl. Mech.*
 895 *Rev.*, 59(1–6), 51–77.

Calvert, J. G., et al. (1978), Mechanism of homogeneous oxidation of sulfur-
 dioxide in troposphere, *Atmos. Environ.*, 12(1–3), 197–226, doi:10.1016/
 0004-6981(78)90201-9.
 Castellanos, P. (2009), Analysis of air quality with numerical simulations
 (CMAQ), and observations of trace gases, PhD dissertation, Univ. of
 Md., College Park.
 Castellanos, P., L. T. Marufu, B. G. Doddridge, B. F. Taubman, J. J.
 Schwab, J. C. Hains, S. H. Ehrman, and R. R. Dickerson (2011), Ozone,
 oxides of nitrogen, and carbon monoxide during pollution events over the
 eastern United States: An evaluation of emissions and vertical mixing,
J. Geophys. Res., 116, D16307, doi:10.1029/2010JD014540.
 CESY (2005), *Energy Statistical Yearbook in 2005*, China Stat. Press,
 Beijing.
 Chin, M., D. J. Jacob, G. M. Gardner, M. S. Foreman-Fowler, P. A. Spiro, and
 D. L. Savoie (1996), A global three-dimensional model of tropospheric sul-
 fate, *J. Geophys. Res.*, 101(D13), 18,667–18,690, doi:10.1029/96JD01221.
 Clarke, J. F., et al. (1997), Dry deposition calculations for the clean air status
 and trends network, *Atmos. Environ.*, 31(21), 3667–3678, doi:10.1016/
 S1352-2310(97)00141-6.
 Community Modeling and Analysis System (2007), CMAQ v4.6 opera-
 tional guidance document, report, Univ. of N. C. at Chapel Hill, Chapel
 Hill.
 Costabile, F., et al. (2006), A preliminary assessment of major air pollutants
 in the city of Suzhou, China, *Atmos. Environ.*, 40(33), 6380–6395,
 doi:10.1016/j.atmosenv.2006.05.056.
 Cox, R. A., and S. A. Penkett (1971), Photo-oxidation of atmospheric SO₂,
Nature, 229(5285), 486–488, doi:10.1038/229486a0.
 Dickerson, R. R., et al. (2007), Aircraft observations of dust and pollutants
 over northeast China: Insight into the meteorological mechanisms of
 transport, *J. Geophys. Res.*, 112, D24S90, doi:10.1029/2007JD008999.
 Dunlea, E. J., et al. (2009), Evolution of Asian aerosols during transpacific
 transport in INTEX-B, *Atmos. Chem. Phys.*, 9(19), 7257–7287, doi:10.5194/acp-9-
 7257-2009.
 Eggleston, A. E. J., and R. A. Cox (1978), Homogeneous oxidation of sulfur-
 compounds in atmosphere, *Atmos. Environ.*, 12(1–3), 227–230, doi:10.1016/
 0004-6981(78)90202-0.
 Ek, M. B., K. E. Mitchell, Y. Lin, E. Rogers, P. Grunmann, V. Koren, G.
 Gayno, and J. D. Tarpley (2003), Implementation of Noah land surface
 model advances in the National Centers for Environmental Prediction
 operational mesoscale Eta model, *J. Geophys. Res.*, 108(D22), 8851,
 doi:10.1029/2002JD003296.
 Fioletov, V. E., et al. (2011), Estimation of SO₂ emissions using OMI retrie-
 vals, *Geophys. Res. Lett.*, 38, L21811, doi:10.1029/2011GL049402.
 Foken, T. (2006), 50 years of the Monin-Obukhov similarity theory, *Bound-
 ary Layer Meteorol.*, 119(3), 431–447, doi:10.1007/s10546-006-9048-6.
 Geng, F. H., et al. (2009), Aircraft measurements of O₃, NO_x, CO, VOCs,
 and SO₂ in the Yangtze River Delta region, *Atmos. Environ.*, 43(3),
 584–593, doi:10.1016/j.atmosenv.2008.10.021.
 Hains, J. C. (2007), A chemical climatology of lower tropospheric trace
 gases and aerosols over the mid-Atlantic region, PhD dissertation, Univ.
 of Md., College Park.
 Hains, J. C., et al. (2008), Origins of chemical pollution derived from Mid-
 Atlantic aircraft profiles using a clustering technique, *Atmos. Environ.*,
 42(8), 1727–1741, doi:10.1016/j.atmosenv.2007.11.052.
 Hand, J. L., and W. C. Malm (2007), Review of aerosol mass scattering
 efficiencies from ground-based measurements since 1990, *J. Geophys.*
Res., 112, D16203, doi:10.1029/2007JD008484.
 Haywood, J., and O. Boucher (2000), Estimates of the direct and indirect
 radiative forcing due to tropospheric aerosols: A review, *Rev. Geophys.*,
 38(4), 513–543, doi:10.1029/1999RG000078.
 He, K. B., et al. (2002), Urban air pollution in China: Current status, character-
 istics, and progress, *Annu. Rev. Energy Environ.*, 27, 397–431, doi:10.1146/
 annurev.energy.27.122001.083421.
 Hong, S.-Y., and J.-O. Lim (2006), The WRF single-moment 6-class micro-
 physics scheme (WSM6), *J. Korean Meteorol. Soc.*, 42, 129–151.
 Hu, H., et al. (2010), Air pollution and control in different areas of China,
Crit. Rev. Environ. Sci. Technol., 40(6), 452–518, doi:10.1080/
 10643380802451946.
 Igarashi, Y., et al. (2006), Seasonal variations in SO₂ plume transport over
 Japan: Observations at the summit of Mt. Fuji from winter to summer,
Atmos. Environ., 40(36), 7018–7033, doi:10.1016/j.atmosenv.2006.06.017.
 Intergovernmental Panel on Climate Change (2007), *Climate Change 2007:
 The Physical Science Basis, Contribution of Working Group I to the
 Fourth Assessment Report (AR4) of the Intergovernmental Panel on Cli-
 mate Change*, 996 pp., Cambridge Univ. Press, Cambridge, U. K.
 Kain, J. S. (2004), The Kain-Fritsch convective parameterization: An update,
J. Appl. Meteorol., 43(1), 170–181, doi:10.1175/1520-0450(2004)043<0170:
 TKCPAU>2.0.CO;2.

- 974 Kan, H. D., et al. (2010), Short-term association between sulfur dioxide and
 975 daily mortality: The Public Health and Air Pollution in Asia (PAPA)
 976 study, *Environ. Res.*, *110*(3), 258–264, doi:10.1016/j.envres.2010.01.006.
- 977 Kim, B. G., et al. (2001), Transport of SO₂ and aerosol over the Yellow sea,
 978 *Atmos. Environ.*, *35*(4), 727–737, doi:10.1016/S1352-2310(00)00344-7.
- 979 Koike, M., et al. (2003), Export of anthropogenic reactive nitrogen and sul-
 980 fur compounds from the East Asia region in spring, *J. Geophys. Res.*, *108*
 981 (D20), 8789, doi:10.1029/2002JD003284.
- 982 Krotkov, N. A., et al. (2006), Band residual difference algorithm for
 983 retrieval of SO₂ from the Aura Ozone Monitoring Instrument (OMI),
 984 *IEEE Trans. Geosci. Remote Sens.*, *44*(5), 1259–1266, doi:10.1109/
 985 TGRS.2005.861932.
- 986 Krotkov, N. A., et al. (2008), Validation of SO₂ retrievals from the Ozone
 987 Monitoring Instrument over NE China, *J. Geophys. Res.*, *113*, D16S40,
 988 doi:10.1029/2007JD008818.
- 989 Larssen, T., et al. (2006), Acid rain in China, *Environ. Sci. Technol.*, *40*(2),
 990 418–425, doi:10.1021/es0626133.
- 991 Lee, C., R. V. Martin, A. van Donkelaar, G. O'Byrne, N. Krotkov, A. Richter,
 992 L. G. Huey, and J. S. Holloway (2009), Retrieval of vertical columns of sul-
 993 fur dioxide from SCIAMACHY and OMI: Air mass factor algorithm devel-
 994 opment, validation, and error analysis, *J. Geophys. Res.*, *114*, D22303,
 995 doi:10.1029/2009JD012123.
- 996 Lee, C., R. V. Martin, A. van Donkelaar, H. Lee, R. R. Dickerson, J. C.
 997 Hains, N. Krotkov, A. Richter, K. Vinnikov, and J. J. Schwab (2011),
 998 SO₂ emissions and lifetimes: Estimates from inverse modeling using in
 999 situ and global, space-based (SCIAMACHY and OMI) observations,
 1000 *J. Geophys. Res.*, *116*, D06304, doi:10.1029/2010JD014758.
- 1001 Li, C., L. T. Marufu, R. R. Dickerson, Z. Li, T. Wen, Y. Wang, P. Wang,
 1002 H. Chen, and J. W. Stehr (2007), In situ measurements of trace gases
 1003 and aerosol optical properties at a rural site in northern China during East
 1004 Asian Study of Tropospheric Aerosols: An International Regional Exper-
 1005 iment 2005, *J. Geophys. Res.*, *112*, D22S04, doi:10.1029/2006JD007592.
- 1006 Li, Z. Q., et al. (2007), Preface to special section on east Asian studies of tro-
 1007 spheric aerosols: An international regional experiment (EAST-AIRE),
 1008 *J. Geophys. Res.*, *112*, D22S00, doi:10.1029/2007JD008853.
- 1009 Li, C., N. A. Krotkov, R. R. Dickerson, Z. Li, K. Yang, and M. Chin
 1010 (2010a), Transport and evolution of a pollution plume from northern
 1011 China: A satellite-based case study, *J. Geophys. Res.*, *115*, D00K03,
 1012 doi:10.1029/2009JD012245.
- 1013 Li, C., Q. Zhang, N. A. Krotkov, D. G. Streets, K. He, S.-C. Tsay, and J. F.
 1014 Gleason (2010b), Recent large reduction in sulfur dioxide emissions from
 1015 Chinese power plants observed by the Ozone Monitoring Instrument,
 1016 *Geophys. Res. Lett.*, *37*, L08807, doi:10.1029/2010GL042594.
- 1017 Li, Z. Q., et al. (2011), East Asian Studies of Tropospheric Aerosols and
 1018 their Impact on Regional Climate (EAST-AIRC): An overview, *J. Geo-
 1019 phys. Res.*, *116*, D00K34, doi:10.1029/2010JD015257.
- 1020 Lin, M. Y., et al. (2008), Long-range transport of acidifying substances in
 1021 East Asia - Part I - Model evaluation and sensitivity studies, *Atmos. Envi-
 1022 ron.*, *42*(24), 5939–5955, doi:10.1016/j.atmosenv.2008.04.008.
- 1023 Liu, X. H., et al. (2010), Understanding of regional air pollution over China
 1024 using CMAQ, part I, performance evaluation and seasonal variation, *Atmos.
 1025 Environ.*, *44*(20), 2415–2426, doi:10.1016/j.atmosenv.2010.03.035.
- 1026 Loughner, C. (2011), Impacts of fair-weather cumulus clouds, bay breezes,
 1027 and landuse on urban air quality and climate, PhD dissertation, Univ. of
 1028 Md., College Park.
- 1029 Loughner, C. P., et al. (2011), Impact of fair-weather cumulus clouds and the
 1030 Chesapeake Bay breeze on pollutant transport and transformation, *Atmos.
 1031 Environ.*, *45*(24), 4060–4072, doi:10.1016/j.atmosenv.2011.04.003.
- 1032 Lu, Z., et al. (2010), Sulfur dioxide emissions in China and sulfur trends in East
 1033 Asia since 2000, *Atmos. Chem. Phys.*, *10*(13), 6311–6331, doi:10.5194/acp-
 1034 10-6311-2010.
- 1035 Luke, W. T. (1997), Evaluation of a commercial pulsed fluorescence detec-
 1036 tor for the measurement of low-level SO₂ concentrations during the gas-
 1037 phase sulfur intercomparison experiment, *J. Geophys. Res.*, *102*(D13),
 1038 16,255–16,265, doi:10.1029/96JD03347.
- 1039 Meng, Z. Y., et al. (2010), Ambient sulfur dioxide, nitrogen dioxide, and ammo-
 1040 nia at ten background and rural sites in China during 2007–2008, *Atmos. Envi-
 1041 ron.*, *44*(21–22), 2625–2631, doi:10.1016/j.atmosenv.2010.04.008.
- 1042 National Center for Atmospheric Research (2010), *Weather Research and
 1043 Forecasting, ARW Version 3 Modeling System's User Guide*, Mesoscale
 1044 and Microscale Meteorol. Div., Boulder, Colo.
- 1045 Okuda, T., et al. (2011), The impact of the pollution control measures for the
 1046 2008 Beijing Olympic Games on the chemical composition of aerosols,
 1047 *Atmos. Environ.*, *45*(16), 2789–2794, doi:10.1016/j.atmosenv.2011.01.053.
- 1048 Otte, T. L. (2008), The impact of nudging in the meteorological model for
 1049 retrospective air quality simulations. Part I: Evaluation against national
 1050 observation networks, *J. Appl. Meteorol. Climatol.*, *47*(7), 1853–1867,
 1051 doi:10.1175/2007JAMC1790.1.
- Pfanz, H., et al. (1987), Mesophyll resistances to SO₂ fluxes into leaves,
 1052 *Plant Physiol.*, *85*(4), 922–927, doi:10.1104/pp.85.4.922.
- Prospero, J. M., D. L. Savoie, and R. Arimoto (2003), Long-term record of
 1053 nss-sulfate and nitrate in aerosols on Midway Island, 1981–2000: Evi-
 1054 dence of increased (now decreasing?) anthropogenic emissions from
 1055 Asia, *J. Geophys. Res.*, *108*(D1), 4019, doi:10.1029/2001JD001524.
- Schlesinger, R. B., and F. Cassee (2003), Atmospheric secondary inorganic
 1056 particulate matter: The toxicological perspective as a basis for health
 1057 effects risk assessment, *Inhal. Toxicol.*, *15*(3), 197–235, doi:10.1080/
 1058 08958370304503.
- Shannon, J. D., and D. L. Sisterson (1992), Estimation of S and NO_x-N
 1059 deposition budgets for the United States and Canada, *Water Air Soil Pol-
 1060 lut.*, *63*(3–4), 211–235, doi:10.1007/BF00475491.
- Singh, H. B., et al. (2009), Chemistry and transport of pollution over the Gulf
 1061 of Mexico and the Pacific: Spring 2006 INTEX-B campaign overview
 1062 and first results, *Atmos. Chem. Phys.*, *9*(7), 2301–2318, doi:10.5194/acp-
 1063 9-2301-2009.
- Sorimachi, A., and K. Sakamoto (2007), Laboratory measurement of the dry
 1064 deposition of sulfur dioxide onto northern Chinese soil samples, *Atmos.
 1065 Environ.*, *41*(13), 2862–2869, doi:10.1016/j.atmosenv.2006.10.025.
- Sorimachi, A., et al. (2003), Measurements of sulfur dioxide and ozone dry
 1066 deposition over short vegetation in northern China—A preliminary study,
 1067 *Atmos. Environ.*, *37*(22), 3157–3166, doi:10.1016/S1352-2310(03)00180-8.
- Stier, P., et al. (2007), Aerosol absorption and radiative forcing, *Atmos.
 1068 Chem. Phys.*, *7*(19), 5237–5261, doi:10.5194/acp-7-5237-2007.
- Stockwell, W. R., P. Middleton, J. S. Chang, and X. Tang (1990), The sec-
 1069 ond generation Regional Acid Deposition Model chemical mechanism for
 1070 regional air-quality modeling, *J. Geophys. Res.*, *95*(D10), 16,343–16,367,
 1071 doi:10.1029/JD095iD10p16343.
- Sun, Y., et al. (2009), Measurement of the vertical profile of atmospheric
 1072 SO₂ during the heating period in Beijing on days of high air pollution,
 1073 *Atmos. Environ.*, *43*(2), 468–472, doi:10.1016/j.atmosenv.2008.09.057.
- Tan, Q., Y. Huang, and W. L. Chameides (2002), Budget and export of
 1074 anthropogenic SO₂ from East Asia during continental outflow conditions,
 1075 *J. Geophys. Res.*, *107*(D13), 4167, doi:10.1029/2001JD000769.
- Taubman, B. F., J. C. Hains, A. M. Thompson, L. T. Marufu, B. G. Doddridge,
 1076 J. W. Stehr, C. A. Piety, and R. R. Dickerson (2006), Aircraft vertical pro-
 1077 files of trace gas and aerosol pollution over the mid-Atlantic United States:
 1078 Statistics and meteorological cluster analysis, *J. Geophys. Res.*, *111*,
 1079 D10S07, doi:10.1029/2005JD006196.
- Thompson, G. E., et al. (2006), A new bulk microphysical parameterization
 1080 for WRF (& MM5), paper presented at Seventh Weather Research and
 1081 Forecasting User's Workshop, NCAR, Boulder, Colo.
- Tu, F. H., D. C. Thornton, A. R. Bandy, G. R. Carmichael, Y. Tang, K. L.
 1082 Thornhill, G. W. Sachse, and D. R. Blake (2004), Long-range transport of
 1083 sulfur dioxide in the central Pacific, *J. Geophys. Res.*, *109*, D15S08,
 1084 doi:10.1029/2003JD004309.
- Twomey, S. (1977), Influence of pollution on shortwave albedo of clouds, *J.
 1085 Atmos. Sci.*, *34*(7), 1149–1152, doi:10.1175/1520-0469(1977)034<1149:
 1086 TIOPOT>2.0.CO;2.
- U.S. Environmental Protection Agency (2004), Air quality criteria for
 1087 particulate matter (final report, Oct 2004), *EPA 600/P-99/002aF-bF*,
 1088 Washington, D. C.
- van Donkelaar, A., et al. (2008), Analysis of aircraft and satellite measurements
 1089 from the Intercontinental Chemical Transport Experiment (INTEX-B) to
 1090 quantify long-range transport of East Asian sulfur to Canada, *Atmos. Chem.
 1091 Phys.*, *8*(11), 2999–3014, doi:10.5194/acp-8-2999-2008.
- Wang, L. T., et al. (2010a), Assessment of air quality benefits from national
 1092 air pollution control policies in China. Part I: Background, emission scenar-
 1093 ios and evaluation of meteorological predictions, *Atmos. Environ.*, *44*(28),
 1094 3442–3448, doi:10.1016/j.atmosenv.2010.05.051.
- Wang, L. T., et al. (2010b), Assessment of air quality benefits from national
 1095 air pollution control policies in China. Part II: Evaluation of air quality
 1096 predictions and air quality benefits assessment, *Atmos. Environ.*, *44*(28),
 1097 3449–3457, doi:10.1016/j.atmosenv.2010.05.058.
- Wang, W., et al. (2008), Aircraft measurements of gaseous pollutants and par-
 1098 ticulate matter over Pearl River Delta in China, *Atmos. Environ.*, *42*(25),
 1099 6187–6202, doi:10.1016/j.atmosenv.2008.06.001.
- Wesely, M. L., and B. B. Hicks (2000), A review of the current status of
 1100 knowledge on dry deposition, *Atmos. Environ.*, *34*(12–14), 2261–2282,
 1101 doi:10.1016/S1352-2310(99)00467-7.
- Witte, J. C., M. R. Schoeberl, A. R. Douglass, J. F. Gleason, N. A. Krotkov,
 1102 J. C. Gille, K. E. Pickering, and N. Livesey (2009), Satellite observations
 1103 of changes in air quality during the 2008 Beijing Olympics and Paralymp-
 1104 ics, *Geophys. Res. Lett.*, *36*, L17803, doi:10.1029/2009GL039236.
- Xue, L. K., et al. (2010), Aircraft measurements of the vertical distribution
 1105 of sulfur dioxide and aerosol scattering coefficient in China, *Atmos. Envi-
 1106 ron.*, *44*(2), 278–282, doi:10.1016/j.atmosenv.2009.10.026.

- 1130 Yang, K., N. A. Krotkov, A. J. Krueger, S. A. Carn, P. K. Bhartia, and P. F.
1131 Levelt (2009a), Improving retrieval of volcanic sulfur dioxide from back-
1132 scattered UV satellite observations, *Geophys. Res. Lett.*, *36*, L03102,
1133 doi:10.1029/2008GL036036.
- 1134 Yang, K., X. Liu, N. A. Krotkov, A. J. Krueger, and S. A. Carn (2009b),
1135 Estimating the altitude of volcanic sulfur dioxide plumes from space-
1136 borne hyper-spectral UV measurements, *Geophys. Res. Lett.*, *36*,
1137 L10803, doi:10.1029/2009GL038025.
- 1138 Yang, K., X. Liu, P. K. Bhartia, N. A. Krotkov, S. A. Carn, E. J. Hughes,
1139 A. J. Krueger, R. J. D. Spurr, and S. G. Trahan (2010), Direct retrieval of
1140 sulfur dioxide amount and altitude from spaceborne hyperspectral UV
1141 measurements: Theory and application, *J. Geophys. Res.*, *115*,
1142 D00L09, doi:10.1029/2010JD013982.
- 1143 Zhang, Q., et al. (2009), Asian emissions in 2006 for the NASA INTEX-B
1144 mission, *Atmos. Chem. Phys.*, *9*(14), 5131–5153, doi:10.5194/acp-9-
1145 5131-2009.
- Zhang, Y. H., et al. (2008), Regional Integrated Experiments on Air Quality 1146
over Pearl River Delta 2004 (PRIDE-PRD2004), Overview, *Atmos. Envi-* 1147
ron., *42*(25), 6157–6173, doi:10.1016/j.atmosenv.2008.03.025. 1148
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- X. Bao and G. Zhao, Henan Meteorological Bureau, Zhengzhou 450003, 1149
China. 1150
- H. He and R. Dickerson, Department of Atmospheric and Oceanic 1151
Science, University of Maryland, College Park, MD 20742, USA. 1152
(hhe@atmos.umd.edu) 1153
- N. A. Krotkov and K. Yang, NASA Goddard Space Flight Center, 1154
Greenbelt, MD 20771, USA. 1155
- C. Li, Z. Li, and C. P. Loughner, Earth System Science Interdisciplinary 1156
Center, University of Maryland, College Park, MD 20740, USA. 1157
- L. Wang and Y. Zheng, Nanjing University of Information Science and 1158
Technology, Nanjing 210044, China. 1159

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